

Second Annual General Meeting, June 12-13, 2019

Integrated Modelling Program for Canada (IMPC), Global Water Futures



Soil Moisture Accounting for Nelson Churchill River Basin using HYPE

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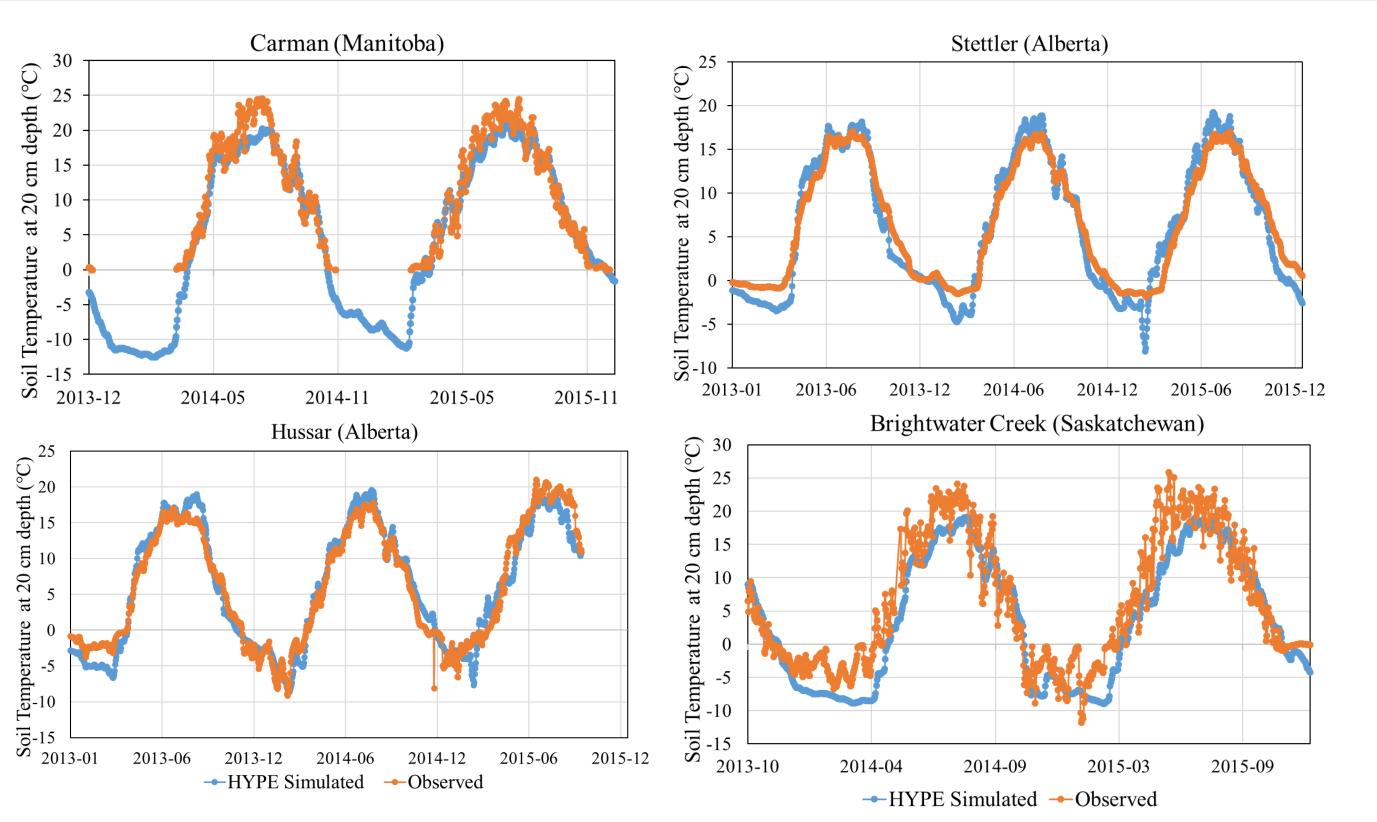
01. Introduction

Hydrological models are important tools to analyse the cold regions processes, such as permafrost, seasonally frozen soil and snow cover, which are widely distributed across Canada. Improvement of the hydrological models to better represent the cold regions processes is one the core objective under Theme A2 of IMPC. In cold regions, frozen soil processes play a key role in generating the runoff by restricting the infiltration during the frozen state, and thawing during the melting phase. Therefore, the improvement of such processes can significantly boost the confidence in model projection for historical and climate change scenarios, and hence reduce the uncertainty associated with it.

Objectives:

- To improve the reliability of HYPE model to simulate soil moisture for NCRB, emphasizing on the frozen soil processes.
- To validate model results with observed soil temperature and soil moisture recorded at





several locations within NCRB

To investigate future climate change impacts on soil moisture and soil temperature using the improved HYPE-NCRB model, and their implications on the uncertainty associated with the projection of streamflow.

NCRB drains over 1.4 millions sq. km

Recorded soil moisture data are available

Most stations record soil temperature and

soil moisture at 20 cm depth, 50 cm depth

meteorological

dataset is used for model setup (Berg et

reanalysis

and 100 cm depth from ground level.

approximately (Fig 1).

starting from 2013.

Hydro-GFD

al., 2018).

02. Study Area

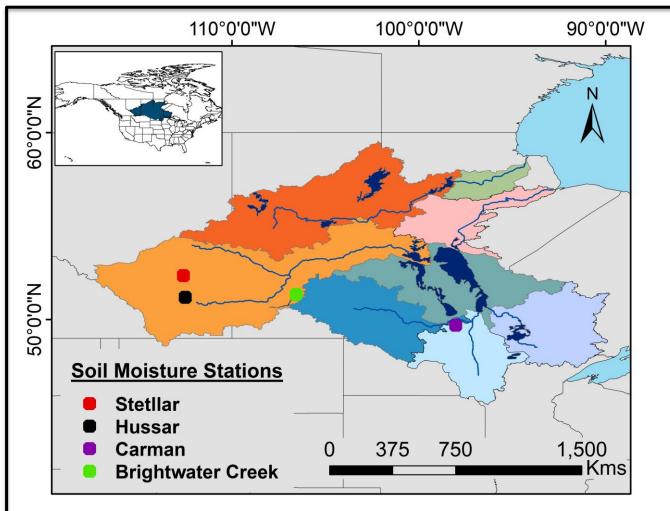


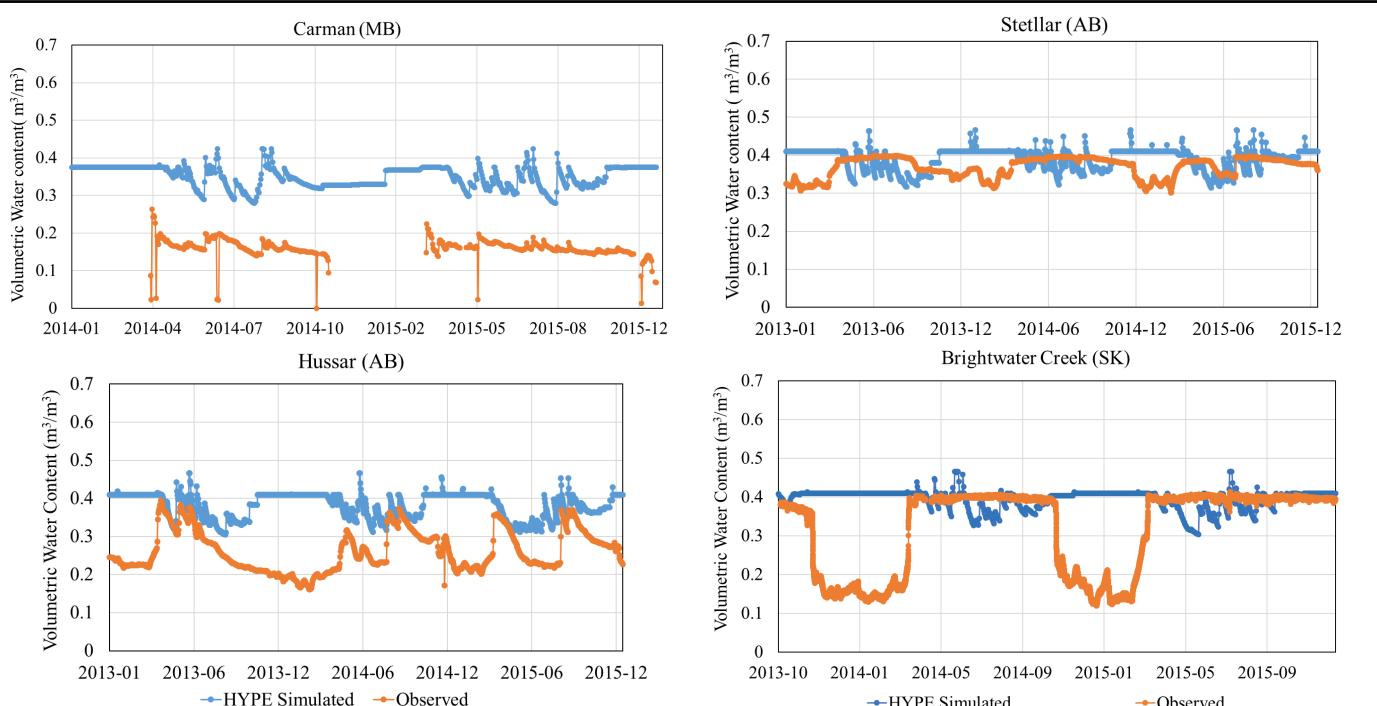
Fig 1 Study area with location of soil moisture monitoring stations within NCRB

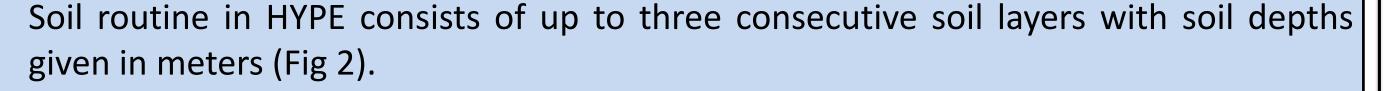
03. Soil Moisture Accounting in HYPE

- HYPE is a semi-distributed model discretized at sub-basin level.
- A sub-basin in HYPE is further divided into classes or SLCs (unique LULC Soil combination), similar to HRUs (Hydrological Response Units).

Fig 3 Simulated soil temperature vs observed soil temperature at several locations within NCRB at 20 cm depth.

- The simulated soil temperature from HYPE (20 cm depth) at different locations were validated with measured temperature (Fig 3).
- Soil temperature simulated in HYPE from 2013 to 2015, at 20 cm depth, matches well with the observed data for all the locations within NCRB.





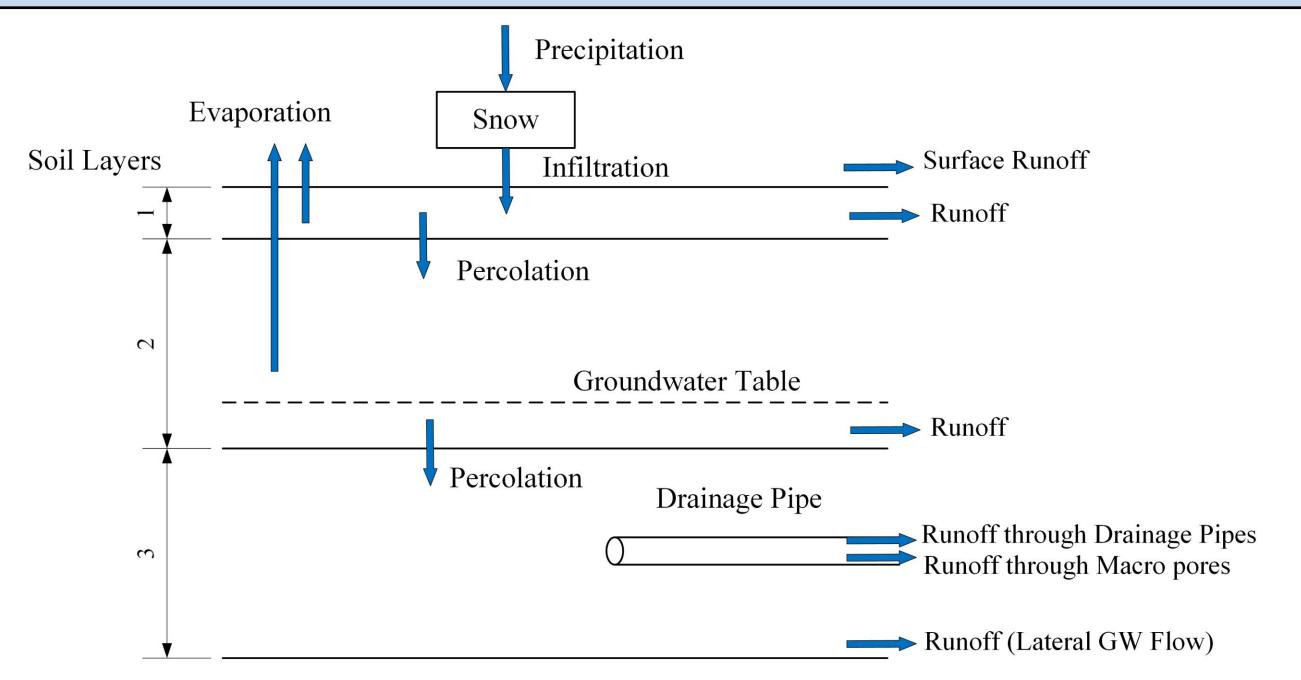


Fig 2 Illustration of flow path in the soil in the HYPE model (Adapted from HYPE wiki page)

- Infiltration in HYPE is calculated from the sum of rainfall and snowmelt.
- If \sum (rainfall+snowmelt) > infiltration capacity, a part of water will not infiltrate into the soil.
- The calculation of actual infiltration considers the effects of surface runoff, macro pore flow and frozen soil.

04. Frozen Soil Infiltration

-HYPE Simulated Observed

Fig 4 Simulated soil moisture vs observed soil moisture at several locations within NCRB at 20 cm depth.

- The soil moisture simulated by HYPE at 20 cm depth fails to capture the soil moisture during the winter months, when infiltration is low and soil moisture drops significantly (Fig 4).
- The simulated soil moisture is significantly higher and remains more or less constant during the freezing period.
- The results could be improved by integrating suitable frozen soil infiltration to limit the soil moisture during winter season, and further calibration of soil moisture parameters will be required.

06. Future work

- Improve HYPE-NCRB model to simulate the soil moisture more accurately, especially during winter.
- The frozen infiltration will be implemented layer by layer based on soil layer temperature.
- The progressive freezing and thawing of soil layers based on soil layer depth will be accounted.
- Analyze future climate change impacts on soil moisture and soil temperature using the improved HYPE-NCRB model, and analyze the uncertainty associated with it.

07. References & Acknowledgements

Berg, P., Donnelly, C., & Gustafsson, D. (2018). Near-real-time adjusted reanalysis forcing data for hydrology. Hydrology and Earth System Sciences, 22(2), 989-1000. Clapp, R., & Hornberger, G. (1978). Empirical equations for some soil hydraulic properties. Water Resources Research, 14, 601-604. MacDonald, M., T.A Stadnyk, S.J Dery, D Gustafsson, K Isberg, B Arheimer. In revision. Improved hydrologic model representation of landscape-based storage in the Hudson Bay Drainage Basin. Submitted to Hydrol. Process. HYP-17-0803.

- If soil temperature > 0°C, liqfrac(k) =1 (normal infiltration, soil water from all layers is available for runoff)
- If soil temperature < 0 °C, liquid water content is calculated based on freezing point depression equation (Clapp and Hornberger, 1978).

 $\theta_{l,max} = \theta_S \left(\frac{L_F (T_{freeze} - T_{soil})}{g T_{soil} \psi_S} \right)^{1/b}$

where, θ_{Lmax} is the maximum volumetric liquid water content (m³ m⁻³), θ_s is the saturation volumetric water content or porosity ($m^3 m^{-3}$), LF is the latent heat of fusion (3.34 \times 105 J kg⁻¹), T_{freeze} = 273.16 K, T_{soil} is soil layer temperature (K), g is gravitational acceleration (9.81 m s⁻²), ψ_s is soil water potential at saturation (m), and b is the shape coefficient of the soil water potential-moisture curve (MacDonald et al., in revision).

Acknowledgements

Financial support for this research was provided by Manitoba Hydro, Natural Sciences and Engineering Research Council of Canada, and Global Water Futures. We are thankful to Agriculture and Agri-Food Canada, and Ministry of Agriculture and Forestry, Alberta for providing soil moisture data. We would also like to acknowledge Dr. Kevin Shook for giving access to WISKI database. We are also grateful to SMHI for providing the HYPE model and the Hydro-GFD data.